

## **Building confidence in projections of the responses of living marine resources to climate change**

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**Building confidence in projections of the responses of living marine resources to climate change**

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24 **Abstract**

25 The Fifth Assessment Report of the Intergovernmental Panel on Climate Change highlights that  
26 climate change and ocean acidification are challenging the sustainable management of living  
27 marine resources (LMRs). Formal and systematic treatment of uncertainty in existing LMR  
28 projections, however, is lacking. We synthesize knowledge on how to address different sources  
29 of uncertainty by drawing from climate model inter-comparison efforts. We suggest an ensemble  
30 of available models and projections, informed by observations, as a starting point to quantify  
31 uncertainties. Such an ensemble must be paired with analysis of the dominant uncertainties over  
32 different spatial scales, time horizons and metrics. We use two examples, (1) global and regional  
33 projections of Sea Surface Temperature and (2) projection of changes in potential catch of  
34 sablefish (*Anoplopoma fimbria*) in the 21<sup>st</sup> century, to illustrate this ensemble model approach to  
35 explore different types of uncertainties. Further effort should prioritize understanding dominant,  
36 under-sampled dimensions of uncertainty, as well as the strategic collection of observations to  
37 quantify, and ultimately reduce, uncertainties. Our proposed framework will improve our  
38 understanding of future changes in LMR and the resulting risk of impacts to ecosystems and the  
39 societies under changing ocean conditions.

40

41 **Living marine resources projections under climate change**

42 The Fifth Assessment Report of the Intergovernmental Panel on Climate Change (IPCC AR5)  
43 highlights that changes in ocean temperature, oxygen, carbonate system, and other ocean  
44 properties are contributing to the challenges of sustainable ocean management (IPCC, 2014). The  
45 importance of a comprehensive assessment of the impact of climate change on the ocean is  
46 highlighted by two new ocean-specific chapters within the IPCC AR5 Working Group II (WGII)

47 on impacts, adaptation and vulnerability (IPCC, 2014). In relation to living marine resources  
48 (LMR), the IPCC Report concludes with medium to high confidence that marine species have  
49 been shifting their ranges, seasonal activities and periodicities, migration patterns, abundances  
50 and inter-/intra- specific interactions that result in changes in trophodynamics in response to  
51 changing ocean conditions (Pörtner *et al.*, 2014). These changes are projected to lead to altered  
52 patterns of ocean productivity, biodiversity and fisheries catch potential in the 21<sup>st</sup> century  
53 (Kirby and Beaugrand, 2009).

54 One of the advances in assessing the impacts of climate change on LMR in the IPCC AR5 WGII  
55 over previous assessment reports is the wider availability and use of ecosystem model  
56 projections. These quantitative model projections include shifts in net primary productivity, the  
57 distribution of exploited populations and changes in potential fisheries production and ecosystem  
58 structure at local and global scales (Pörtner *et al.*, 2014). Projections have been generated from  
59 modelling approaches that range from global coupled atmosphere-ocean-biogeochemistry earth  
60 system models (e.g., Bopp *et al.*, 2013), to species distribution models (e.g., Cheung *et al.*,  
61 2009), single-species population dynamic models (e.g., Lehodey *et al.*, 2010), and whole  
62 ecosystem models (e.g., Ainsworth *et al.*, 2011; Griffith *et al.*, 2011). The scope, objectives,  
63 assumptions, scales (spatial and temporal) and degree of validation with empirical data vary  
64 widely across these models, and approaches range from highly empirical to highly mechanistic  
65 (Barange *et al.*, 2010; Fulton, 2010; Plagányi *et al.*, 2011; Stock *et al.*, 2011).

66 Statements of confidence concerning the impacts of climate change on LMRs within the IPCC-  
67 AR5 WGII report were based on a qualitative assessment of observational evidence and  
68 individually published projections encompassing the diversity of LMR models described above.  
69 While this is a necessary starting point, more quantitative confidence estimates for projections

70 can increase their utility for policy formulation and evaluation. There is therefore a need for a  
71 quantitative framework for systematically exploring uncertainties in LMR projections. Such a  
72 framework would also help identify where investment in further theoretical development,  
73 observational measurements, and model development are needed, ultimately improving the  
74 reliability of climate-LMR projections (Cheung *et al.*, 2013a; Brander, 2015). Systematic  
75 exploration of uncertainties have been undertaken for climate and oceanographic projections  
76 (e.g., the Atmospheric Model Intercomparison Project (Gates, 1992) and the Coupled Model  
77 Intercomparison Project (Meehl *et al.*, 2000; Taylor *et al.*, 2011) and for impact assessments of  
78 selected sectors (e.g., Agricultural Model Intercomparison and Improvement Project  
79 (Rosenzweig *et al.*, 2013)). Exploration of uncertainties are also an important component in  
80 traditional fisheries resource assessment, while the increasing demand for ecosystem-based  
81 fisheries management raises additional challenges to systematically understanding projection  
82 uncertainties (e.g., Hill *et al.*, 2007; Link *et al.*, 2012). More recently, initiatives on comparing  
83 fisheries models (e.g., Fisheries Model Intercomparison Project, ICES-PICES Strategic Initiative  
84 on Climate Change Impacts on Marine Ecosystems) have also been started.

85 While challenges in quantifying uncertainty in climate-LMR projections for global change  
86 assessment parallel those considered in modelling other complex natural systems such as  
87 climate, there are additional sets of complexity that are specific to LMRs. Climate-LMR  
88 projections require linking physical, biological and human sub-systems across different temporal  
89 and spatial scales. Such inter-linkages lead to additional uncertainties that originate from  
90 particular systems or scales (Planque, 2015). In addition, the behavior of some components of  
91 LMR systems is difficult to predict, such as the responses of fishing activities to changes in  
92 climate and fisheries resources). Moreover, many LMR models require large number of input

93 parameters relative to the available observational data that are available to calibrate and validate  
94 the model outputs. Techniques for assessing model uncertainties that are commonly applied to  
95 conventional fisheries assessment (e.g., Bayesian estimates of process and observation errors) are  
96 thus difficult to apply to climate-LMR projections.

97 This paper aims to synthesize our knowledge of the uncertainties of LMR projections under  
98 climate change and propose a framework to systematically assess such uncertainty. Our paper  
99 complements that of Payne *et al.* (this volume), which reviews existing approaches in addressing  
100 uncertainties in LMR. Here, we focus on the following: firstly, we characterize different types of  
101 uncertainty in climate and LMR projections, highlighting the challenges of the large uncertainty  
102 space; secondly, drawing from the experience of physical climate model inter-comparisons, we  
103 explore how multi-model comparison and ensemble frameworks can be used to systematically  
104 identify and quantify uncertainties in LMR projections. Through an example, we highlight the  
105 relative roles of uncertainty linked to climate variability, climate model uncertainty and future  
106 emissions scenarios as a function of time horizon and spatial scale. This is followed by a  
107 discussion of the role of observations in refining uncertainty estimates. Finally, we discuss how  
108 outcomes from this model-assessment framework can be used to evaluate the risk of climate  
109 change to LMRs and inform the design of management and conservation measures to reduce  
110 such risk.

111

## 112 **Sources of uncertainty**

113 Climate-LMR models that estimate the impacts of climate change generally have three model  
114 components that are linked to describe the responses of marine resources, fisheries, and human

115 society to climate systems. These components generally include an atmosphere-ocean-  
116 biogeochemical and lower-trophic level models, a fish or upper-trophic level model (Holt *et al.*,  
117 2014), and a model for the extraction and availability of ecosystem services from marine  
118 ecosystems (see Fulton, 2010; Plagányi *et al.*, 2011; Stock *et al.*, 2011). The three components  
119 are either related “off-line”, where each model component is run separately with the outputs  
120 from one component used as inputs for another (Cheung *et al.*, 2011; Blanchard *et al.*, 2012;  
121 Christensen *et al.*, 2015), or dynamically (i.e., "on-line") with the models incorporating fully  
122 interactive processes and, in some cases, feedbacks among the three components (Fulton, 2010;  
123 Lefort *et al.*, 2015).

124 Research on physical climate projections, biodiversity and ecological modelling has recognized  
125 numerous topologies of uncertainties (Regan *et al.*, 2002; Link *et al.*, 2012). Modelling of  
126 physical and biogeochemical properties of atmospheric and ocean systems in climate change  
127 assessments have commonly categorized uncertainties, for any time horizon and spatial scale,  
128 into three components: (1) internal variability, (2) model uncertainty, and (3) scenario  
129 uncertainty (Table 1) (Hawkins and Sutton, 2009). In our discussion of the uncertainties  
130 associated with climate-LMR projections, we adopt this terminology to leverage the knowledge  
131 and experience of the climate modelling communities.

132 Internal variability is caused by natural physical and ecological processes that are intrinsic to  
133 climate and ecological systems. It arises in both temporal and spatial dimensions, even in the  
134 absence of any external (e.g., anthropogenic) perturbations and includes phenomena such as the  
135 El Niño Southern Oscillation (ENSO), the North Atlantic Oscillation (NAO), the Atlantic  
136 Multidecadal Oscillation (AMO), variations in gyre boundaries not correlated to major climate  
137 models, and predator-prey cycles etc. (Day, 1982). Century-scale climate change projections

138 developed in association with the IPCC realistically resolve many modes of internal climate  
139 variability, but these simulations are not designed to simulate a specific observed event or predict  
140 a future event, and will not capture all aspects of spatial and temporal scales of these modes  
141 (Guilyardi *et al.*, 2009). For ecological systems, natural fluctuations that are driven by  
142 environmental variability and dynamics of ecological interactions are often difficult to predict  
143 (Beckage *et al.*, 2011; Deser *et al.*, 2012), causing systematic or seemingly random variations in  
144 ecological states that may persist for a decade or more (Deser, 2013; Stocker *et al.*, 2014).  
145 Different initial conditions of the climate or LMR models, representing different realizations of  
146 the climatic and ecological systems, will generate different patterns of internal variability. Thus,  
147 one method to explore internal variability is to analyze simulation results generated from  
148 ensemble members of climate and ecological models that have different initial conditions.

149 Model uncertainty is comprised of two sub-categories: parameter and structural uncertainty  
150 (Tebaldi and Knutti, 2007). Parameter uncertainty relates to the specific parameter values used in  
151 the formulae that influence the behavior of a model (Tebaldi and Knutti, 2007; Knutti *et al.*,  
152 2010). For parameters that are estimated from observations, parameter uncertainty stems from  
153 our limited ability to precisely measure or estimate specific physical or ecological processes and  
154 quantities (Link *et al.*, 2012), as well as from the inherent variability in certain processes (e.g.  
155 growth rates that vary across individuals) that are not resolved within the models.

156 Structural uncertainty relates to the spatial, temporal, and mathematical resolution employed by a  
157 model and the types of processes that are represented. Structural uncertainty includes the  
158 function forms of equations used to describe mechanistic processes and the types of interactions  
159 assumed to influence climate-LMR processes. Such uncertainties cannot be explored via  
160 parameter perturbations. For example, explicit trophic relationships that are not described by



161 size-structured interactions are not represented in size-based trophodynamic models (e.g.,  
162 Blanchard *et al.*, 2012; Watson *et al.*, 2014), while such relationships may be included in  
163 functional-group type food web models (e.g., Christensen and Walters, 2004).

164 Scenario uncertainty relates to the many possible futures comprising different socio-economic  
165 policies and technological developments likely occurring over the course of a model projection  
166 (e.g., Moss *et al.*, 2010; Nakicenovic *et al.*, 2014). Climate-LMR model drivers include the  
167 spatial and temporal changes in greenhouse gas and aerosol concentrations, fishing effort, and  
168 other human social-economic activities. Scenario uncertainty is not completely independent of  
169 internal variability in the climate-LMR system, as future decisions on the utilization and  
170 conservation of resources are sensitive to natural variation in the availability and distribution of  
171 LMR (e.g., the fishing quota decided on for the next management cycle are dependent on the  
172 productivity and abundance of the resources, as well as on how neighboring countries or regions  
173 are managing their resources).

174 The full range of possible future states for a given LMR reflects contributions from all of the  
175 sources of uncertainty outlined above, with potential cascades of uncertainties interacting and  
176 accumulating over components of the climate-LMR models (Figure 1). For any particular  
177 scenario, LMR models that differ in their structure and parameter values will simulate a range of  
178 future changes in ocean biogeochemistry, fish and fisheries. Additionally, an individual model  
179 with a fixed set of parameters will display variability in projections as a result of the internal  
180 variability associated with natural fluctuations of the climatic or ecological systems.  
181 Uncertainties that originate from different climate-LMR model sub-components may be additive  
182 or multiplicative. Thus, the final scope of uncertainties of LMR projections is expected to be  
183 different from the uncertainty scope of each model sub-component.

[Figure 1. ]

184

185 The width of the envelope of uncertainty is dependent on the nature of interactions between  
186 linked models; the types of interactions include linearity of the linkages, existence of threshold  
187 responses, and positive/negative feedbacks (Peters and Herrick, 2004). When the processes  
188 linking two or more models are non-linear, uncertainties may be dampened or magnified  
189 through model linkages, for example, through attenuation or amplification of changes in higher  
190 trophic level production in marine ecosystems driven by climate change (Chust *et al.*, 2014;  
191 Stock *et al.*, 2014a). Feedbacks in social-ecological systems can be positive or negative, and  
192 uncertainties propagated in models that are linked dynamically with feedbacks resulting in  
193 emergent dynamics are difficult to predict.

194

195 Here, we draw experience from the large body of research on exploring uncertainties of climate  
196 projections to propose that the envelope of uncertainties of climate-LMR projection can be  
197 explored by systematically quantifying the three categories of uncertainty that we discussed  
198 above: internal variability, model uncertainty and scenario uncertainty. Review on specific  
199 techniques to explore each source of uncertainties can be found in Payne *et al.* (this volume).

200

### 201 **Experiences from quantifying uncertainty of climate projections**

202 For ocean-atmospheric general circulation models and biogeochemical models, the Coupled  
203 Model Intercomparison Project Phase 5 (CMIP5) multi-model database allows assessment of  
204 uncertainty in climate change projections across the dimensions illustrated in Fig. 1. Climate

205 change projections were produced from more than 30 models, developed by different modelling  
206 groups with a standard set of scenario experiments (Flato *et al.*, 2013). The CMIP5 database  
207 allows some exploration of uncertainty, but comprehensive categorization of uncertainty into  
208 structural uncertainty, parameter uncertainty, and internal variability is not possible. The main  
209 challenges include the limited number of modeling groups that were able to contribute ensembles  
210 of runs, some models are fully independent of one another, and a lack of exploration of  
211 parameter uncertainty. Ideally, the ensemble should consist of a random sample across the  
212 uncertainty components in Fig. 1. For complex inter-linked models such as climate or climate-  
213 LMR models, exploring their full scope of uncertainty would require substantial computational  
214 time and other resources. Thus, a systematic approach is needed to efficiently explore the  
215 envelope of uncertainties.

216 To further explore the uncertainty contributed by internal variability for each model, ensembles  
217 of climate simulations have been run under identical forcing, but with each simulation initialized  
218 with slightly different, but equally plausible, conditions (Rodgers *et al.*, 2015). The chaotic  
219 nature of climate variability quickly produces different climate trajectories in each ensemble  
220 member (Wittenberg *et al.*, 2014). By considering each of the trajectories as a plausible outcome,  
221 the ensemble can be used to isolate that part of projection uncertainty due to internal variability  
222 (Frölicher *et al.*, 2009; Deser *et al.*, 2012).

223 Hawkins and Sutton (Hawkins and Sutton, 2009) analyze CMIP3 (i.e., the precursor of CMIP5)  
224 projections to explore the contribution of internal variability and model and scenario  
225 uncertainties to climate projections at global and regional scales. They showed that the dominant  
226 sources of uncertainty in surface air temperature projections vary with spatial scale and time  
227 horizon, noting the importance of model uncertainty and internal variation for mid-21st century

228 regional projections. To further illustrate the application of the framework used by Hawkins and  
229 Sutton (2009) in the oceanic realm, we analyzed the projection uncertainties for sea surface  
230 temperature by combining CMIP5 projections and a large ensemble projections from the Earth  
231 System Model of the Geophysical Fluid Dynamic Laboratory (GFDL ESM2M model; Dunne *et*  
232 *al.*, 2012; Dunne *et al.*, 2013; Rodgers *et al.*, 2015).

233

234 We used the projection of SST as an example of exploring the sensitivity of model projections to  
235 different sources of uncertainties. Scenario uncertainty is estimated to be the difference between  
236 the multi-model mean of projections from 15 CMIP5 models of two 21st century emissions  
237 scenarios: the low-emissions scenario RCP2.6 with an increased radiative forcing that peaks at  
238 approximately 3 W/m<sup>2</sup> before 2100 and then declines to 2.6 W/m<sup>2</sup> by 2100, and the high-  
239 emissions scenario RCP8.5, with an increased radiative forcing of >8.5 W/m<sup>2</sup> by year 2100  
240 (Meinshausen *et al.*, 2011). Model uncertainty is estimated as the standard deviation of changes  
241 in SST (10-year running mean) from each model projections. The internal variability is estimated  
242 as the standard deviation of projections from 30 ensemble member simulations of GFDL  
243 ESM2M (Rodgers *et al.*, 2015).

244

245 [Figure 2.]

246 Globally, the analysis shows that model uncertainty is dominant in the medium term SST  
247 projection (2030 - 2050), while the long-term (2080 - 2100) projection is dominated by scenario  
248 uncertainty (Fig. 2). The large model uncertainty over the medium term reflects the large  
249 variations in regional scale biases in the models. Although the importance of internal variability

250 is second to model uncertainty in near term projection (2010 – 2030), its relative importance  
251 decreases rapidly further into the future.

252 The relative importance of different uncertainty sources varies between different regions. In the  
253 Northeast Atlantic (North Sea Large Marine Ecosystem, (Pauly *et al.*, 2008)), the importance of  
254 scenario uncertainty is smaller compared to those projections at the global scale, while model  
255 uncertainties and internal variability become the dominant uncertainty sources. The internal  
256 variability in the Northeast Atlantic may represent known properties of interannual and  
257 multidecadal climate and oceanographic variability such as North Atlantic Oscillation (NAO)  
258 and Atlantic Multidecadal Oscillation (AMO) (Viles and Goudie, 2003; Beaugrand and Kirby,  
259 2010). In the Northeast Pacific (Gulf of Alaska Large Marine Ecosystem), internal variability  
260 becomes a dominant source of uncertainty representing properties, such as ENSO and Pacific  
261 Decadal Oscillation (PDO). In both basin scale examples, the internal variability of SST is a  
262 bigger contribution to projection uncertainty than in the global scale projection (Fig. 3).  
263 Moreover, in the short to medium term, the projected increase in SST is not sensitive to different  
264 emission scenarios, both globally and in the NE Atlantic and NE Pacific (Fig. 3). However, long-  
265 term warming is much more sensitive to different emission scenarios, particularly at the global  
266 scale. We anticipate the increased importance of internal variability observed at the basin scale  
267 may be even more prominent when examining even smaller spatial scales.

268 [Figure 3.]

269 In addition to highlighting the relative contribution of different sources of uncertainty, this  
270 exploration of uncertainty suggests strategies to prioritize investment in order to improve our  
271 understanding of specific types of uncertainties. In the example presented here, the large model

272 uncertainties in the projections of SST in the Northeast Atlantic call for better understanding of  
273 key processes that may be represented differently among models. In the Northeast Pacific, where  
274 large internal variability is difficult to reliably predict, the medium term effects of greenhouse  
275 gas emission will be difficult to separate from natural variability. This further highlights the need  
276 for better understand inter-annual variability and thus the need for longer-term observational  
277 records.

278

### 279 **Systematic exploration of climate-LMR projection uncertainties**

280 Systematic exploration of the components of uncertainty in both space and time dimensions in a  
281 manner analogous to examples from physical climate model projections (Figure 3) is critical for  
282 moving quickly toward refined uncertainty bounds on climate-LMR projections. Thus,  
283 exploration of uncertainties within climate-LMR projections would include: (1) making  
284 projections from ensemble members of models with different properties of internal temporal or  
285 spatial variability; (2) making projections from ensemble members of models with different  
286 model structure and parameter values, and (3) generating projections that are based on different  
287 climate and fishing scenarios.

288 The conditions to systematically explore uncertainties within climate-LMR projections already  
289 exist. For fish and fisheries models, attempts to explore the full matrix of uncertainties  
290 (particularly model uncertainty with scenario uncertainty) have been made for a limited number  
291 of fisheries or stocks (Table 2). Existing examples mainly involve Management System  
292 Evaluations in which the performance of different models is assessed under different  
293 management scenarios (Link *et al.*, 2012). Methods such as Monte Carlo simulation, Bayesian

294 statistical frameworks, and a plethora of quantitative methods also provide a basis for exploring  
295 both the parameter and structural components of model uncertainty (Hill *et al.*, 2007; Hollowed  
296 *et al.*, 2013). Moreover, various statistical approaches are available to analyze the properties of  
297 different components of uncertainty, and how they contribute to the full scope of uncertainty  
298 (Saltelli *et al.*, 2000). Furthermore, initiatives such as the fisheries component of the Inter-  
299 Sectoral Impact Model Intercomparison Project (ISI-MIP) (Warszawski *et al.*, 2014), which aim  
300 to develop LMR projection databases for climate-fisheries assessment that are similar in nature  
301 to CMIP now been established. Such a database would facilitate collaborative efforts of LMR  
302 research communities to explore the full scope of uncertainties.

303 A remaining knowledge gap in climate-LMR uncertainty exploration is the limited  
304 understanding of uncertainties arising from internal variations in the ecological system or fishing  
305 scenarios in projecting LMR changes, as well as their interactions with internal variability at  
306 different temporal and spatial scales. The linkages between physical and biogeochemical ocean  
307 changes and ecosystem responses are likely to be non-linear and may also involve thresholds;  
308 thus the resulting pattern of internal variability of climate-LMR model projections are likely to  
309 be more complex. For example, the actual response of LMRs to a particular level of  
310 environmental change may be limited by predator – prey interactions, or altered by species-  
311 specific sensitivity and adaptability to environmental fluctuations (Foden *et al.*, 2013).

312 Exploration of internal variability in climate-LMR projections can be done by comparing  
313 projections from ensemble members of a single model with different sets of initial conditions.  
314 For example, we used three versions of Dynamic Bioclimate Envelope Model (DBEM) (Cheung  
315 *et al.*, 2011; Cheung *et al.*, in review) to project changes in maximum potential catch of sablefish  
316 (*Anoplopoma fimbria*) in the Northeast Pacific (Gulf of Alaska Large Marine Ecosystem) from

317 2000 to 2060 (Figure 4). Specifically, we explored the effects of internal variability of ocean  
318 conditions using 20 different ensemble member projections from the GFDL ESM2M (Rodgers *et*  
319 *al.*, 2015). We also compared the relative contribution of uncertainties from internal variability  
320 of ocean conditions, structural uncertainties of DBEM, and uncertainty from different climate  
321 scenarios (RCP 2.6 and 8.5).

322 The results suggest that internal variability is a dominant source of uncertainty for sablefish in  
323 the Northeast Pacific (Gulf of Alaska Large Marine Ecosystem) by 2060 relative to 2000,  
324 followed by the structural uncertainties of DBEM. Scenario uncertainty contributes less than  
325 10% of the total uncertainty. This is broadly consistent with the projected SST changes in this  
326 region, with internal variability contributing around 40% to 70% of the total uncertainty over the  
327 time frame of 2000 - 2060 (Figure 2). However, model uncertainty is substantially lower for  
328 sablefish projections relative to SST projections, possibly because the structural difference  
329 between CMIP5 models (used in SST projection) is much larger than those between the three  
330 versions of DBEM (used in sablefish projection). Also, in addition to SST, DBEM projections  
331 are driven by other ocean biogeochemical variables, such as oxygen and net primary production  
332 (Cheung *et al.*, 2011). Internal variability of multiple oceanographic properties may have  
333 magnified the internal variability of the DBEM projections.

334 [Figure 4.]

335 Since DBEM outputs represents mainly long-term trend of potential catches, inter-annual  
336 variation of reported catches is substantially higher than the internal-variability of the projections  
337 (Fig. 5). DBEM does not represent some processes that contribute to inter-annual variability of  
338 catches such as recruitment variability and changes in fishing effort. Besides spawning stock



339 abundance, recruitment variability could be dependent on both physical (temperature, wind,  
340 current) and/or biological (primary productivity, predation pressure) at different spatial and  
341 temporal scales (Houde, 2008). The relative importance of these factors and the processes  
342 contributing to recruitment vary between species. In addition, catches are also dependent on  
343 changes in fishing effort which can be dependent fisheries management (e.g., quota), social-  
344 economics factors (e.g., price of fish and cost of fishing), and fishers' behavior. DBEM does not  
345 resolve many of these processes and does not have species-specific recruitment sub-model.  
346 Therefore, DBEM is not expected to represent the actual inter-annual variability of the catch. On  
347 the other hand, DBEM is structured to represent the long-term trends of resource productivity.  
348 The long-term trend (20-year running mean) of the reported catch of sablefish falls within the  
349 range of trajectories of the projections (Fig. 5).

350 [Figure 5]

351 The example of the sablefish highlights the need to carefully consider the actual processes that  
352 are represented by the sample of LMR models in quantifying uncertainty from model ensembles.  
353 This challenge applies to both ocean biogeochemical and LMR models. For instance, the  
354 relatively coarse-resolution Earth System Models do not capture potentially large random  
355 variability associated with submesoscale and mesoscale ocean features such as fronts, eddies and  
356 filaments (Stock *et al.* 2011).

357 A standardized set of climate-LMR scenarios is needed to quantify scenario uncertainty for  
358 climate-LMR projections. These scenarios must be reconciled with a range of different  
359 realizations of future emission (e.g., IPCC AR5's Representative Concentration Pathways, or  
360 RCPs) (Moss *et al.*, 2010) and social-economic development (e.g., Shared Socio-economic

361 Pathways or the Sustainable Development Goals) (Griggs *et al.*, 2013; Hunter and O'Neill,  
362 2014). However, emissions scenarios only describe broad-brush societal changes in the 21<sup>st</sup>  
363 century. Fishing sector-specific storylines concerning management, aquaculture and  
364 technological development, and demand for fish in countries across the economic development  
365 spectrum at global and regional scales are also needed. Such factors would ultimately affect the  
366 magnitude and distribution of fishing effort. Trajectories of other human marine-related activities  
367 that drive changes in marine ecosystems should also be included (Figure 6). Development of  
368 these scenarios requires interdisciplinary collaboration between natural and social scientists  
369 (Österblom *et al.*, 2013). Although, there are currently independent efforts to develop such  
370 scenarios at global and regional scales (e.g., Barange *et al.*, 2014; Jones *et al.*, 2014),  
371 community-wide effort in developing standardized sets of scenarios would facilitate consistent  
372 comparison of LMR projections.

373 [Figure 6]

374

### 375 **Building confidence and constraining the scope of plausible projections with observations**

376 Observations across different scales are critical for building confidence in projections and  
377 reducing the scope of LMR uncertainty by constraining parameters, model structures, and  
378 eliminating implausible solutions. Model metrics are observations that can be compared to model  
379 outputs in order to obtain a quantitative assessment of model skill. Model metrics that are of  
380 particular interest for LMR models include species distributions, and the composition and  
381 abundance of fisheries catches (Table 3). These data are generally available for broad-scale  
382 evaluation.

383 Different LMR models may vary in their ability to represent seasonal cycles, inter-annual  
384 variability and/or long-term (decadal or longer) trends, with skill at one scale not always  
385 implying skill at others (Table 3). To assess confidence in the temporal properties of climate-  
386 LMR projections, we suggest three possible tiers of evaluation that involve the use of  
387 observational data to assess consistency with: (1) mean observed spatial patterns or seasonal  
388 climatologies across the scale of interest; (2) previously observed responses to climate  
389 variability; and (3) observed long-term trends attributable to climate change, fishing and other  
390 human drivers. In practice, the ability to assess models across all three tiers is often limited by  
391 data availability, particularly the paucity of the long-term, comprehensive, and high-quality data  
392 sets required to assess models against the often subtle trends in tier 3 (see next section).  
393 Comparisons between LMR data and model projections are also challenging due to issues of  
394 consistency between timeframe and spatial scales, as well as the confounding effects of multiple  
395 human pressures, such as climate and fishing (McOwen *et al.*, 2014). These challenges should  
396 not, however, preclude improving confidence in LMR-climate projections.

397 Confidence in climate-LMR model projections can arise from model evaluation across a subset  
398 of tiers, as well as the reliance of models on robust physiological and ecological principles  
399 (Stock *et al.*, 2011). Real caveats, however, are needed. Observational limitations also suggest  
400 that great care should to be taken eliminating particular projections from consideration within an  
401 ensemble framework. That is, a coarse culling of grossly inconsistent simulations (Overland *et*  
402 *al.*, 2011) is suggested rather than attempting to finely weight models based on nuanced  
403 differences in model-data fit. Even if model projections fit well with observational data, it does  
404 not guarantee that the model can accurately predict future changes, particularly when future  
405 conditions (environmental conditions or human activities) lie outside the bounds of historical

406 conditions. In addition, a good fit between model projections and observational data could, on  
407 occasion, be more indicative of over-parameterization rather than prediction skill.

408

## 409 **Observation data and model metrics**

410 In the paragraphs that follow, we review available LMR data and their potential use as model  
411 metrics for evaluating LMR projections across the three tiers of evaluation described previously.  
412 We focus on the utility of three broad categories of LMR observations: fisheries dependent data,  
413 scientific surveys, and species occurrence records. Similar efforts focusing on metrics for  
414 physical climate models (Knutti and Sedláček, 2013) and biogeochemical/plankton food web  
415 models (Stock *et al.*, 2014b) are also being undertaken. We also identify key uncertainties  
416 associated with such observational data, as these would complicate their use in assessing the  
417 reliability of LMR projections.

### 418 a. Fisheries dependent data

419 Fisheries catch data are particularly useful for Tier 1 and 2 evaluations as they are of direct  
420 relevance to LMRs and their broad spatial, temporal and taxonomic coverage. Total catch  
421 potential can be estimated from the maximum catch of historical time-series, under certain  
422 assumptions concerning fishing effort (Cheung *et al.*, 2008; Friedland *et al.*, 2012). Moreover,  
423 spatial patterns and temporal changes in catch volume (Cheung *et al.*, 2013c) and functional and  
424 taxonomic composition (Cheung *et al.*, 2013b) of fisheries catch can be obtained from global  
425 fisheries databases. Species composition can be aggregated by body size-classes (for size-based  
426 LMR models (Blanchard *et al.*, 2012)), functional role (for functional group trophodynamic  
427 models (Christensen and Walters, 2004)), and by species (for species distribution models

428 (Cheung *et al.*, 2011)). Fisheries catch data can be obtained from the Sea Around Us project  
429 (SAU) ([www.searoundus.org](http://www.searoundus.org)), which provides spatially explicit estimates of global catches  
430 from 1950 onward. In addition, the recent effort of SAU to reconstruct catches that are not  
431 reported in the United Nations Food and Agriculture Organization (FAO) landings statistics  
432 further improves the utility of such data for use as a metric for model comparisons (e.g., Zeller *et al.*  
433 *al.*, 2006). For example, in the Northeast Atlantic, fisheries catch and effort data since the early  
434 20<sup>th</sup> century can be used to understand the ability of LMR models to reproduce changes driven  
435 by the Atlantic Multi-decadal Oscillation and the North Atlantic Oscillation (Kerby *et al.*, 2013).  
436 Similar examples of the potential use of long-term series of fish and fisheries data are also  
437 available in the Northeast Pacific (Lindgren *et al.*, 2013), and large pelagic long-line catch data  
438 are also available for ocean basins. As such datasets are spatially-explicit, the estimated catch-  
439 per-unit-effort can be used as an indicator of the distribution of large pelagic fishes, including  
440 tunas, billfishes, and sharks (Myers and Worm, 2003). Annual and decadal patterns of catches  
441 and their compositions can be assessed to understand the ability of the model to reproduce  
442 interannual and long-term changes in fisheries catches. Interpretation of fisheries catch data,  
443 however, must be done with care as changes or differences in fishing effort, gear, regulations,  
444 taxonomic identification, economics, or human behavior can strongly affect the quantity,  
445 composition, and location of catches (Pinsky and Fogarty, 2012). For this reason, determining  
446 whether observed changes in catch data are caused by climate, ecology, or human behavior can  
447 be complicated. Fisheries dependent data have substantial uncertainties because of inconsistent  
448 data quality and biases in sampling methods, timing and location. Fisheries catches and landings  
449 data may be under-reported (Zeller *et al.*, 2006), over-reported (Watson and Pauly, 2001) or  
450 mis-reported (Pascoe *et al.*, 2001), and the reliability and accuracy of the data may change over

451 time. Also, biases in the location and timing of fishing activities render it challenging to  
452 standardize and use fisheries dependent catch-per-unit-effort data as an index of abundance  
453 (Maunder *et al.*, 2006). There may therefore be biases in using such data to interpret resource  
454 abundance and distribution (Walters, 2003).

455 b. Scientific survey data

456 Scientific surveys are useful across all three tiers of evaluation. They can provide spatial and  
457 temporal patterns of abundance, biomass, biodiversity and distribution. Among the benefits of  
458 scientific surveys is the use of standardized and repeatable methods, stratified random or fixed  
459 design to facilitate statistical inference, and documented survey locations so that both species  
460 presence and absence can be known. These properties make it more likely to attribute observed  
461 changes to particular drivers, such as fishing, pollution and climate change, compared to fisheries  
462 data. For example, data from the California Cooperative Oceanic Fisheries Investigations  
463 (CalCOFI) (Bograd *et al.*, 2003) for the California Current Large Marine Ecosystem, which is  
464 strongly affected by decadal to multidecadal atmospheric oscillations, such as ENSO and PDO,  
465 provide detailed documentation of ecological changes since 1951. The CalCOFI data describe  
466 the abundance of plankton, including larval fishes. A time series of larval fish abundance provide  
467 a useful proxy for adult fish abundance (Koslow *et al.*, 2013). Some surveys further record  
468 information on oceanographic conditions, which might be useful for simultaneously assessing  
469 the skill of the climatic and ecological components of LMR models. Although a number of  
470 surveys available have been sampling for more than four decades, care must be taken to ensure  
471 that large changes in survey methods have not biased the time-series. A common standardization  
472 is to ensure that the same region has been surveyed consistently through time. Also, bias

473 correction factors may be available to account for changes in survey methods (e.g., Ohman and  
474 Smith, 1995).

475 Although survey data can provide estimates of large-scale changes in the distribution of relative  
476 abundance or biomass of LMR (e.g., Pinsky *et al.*, 2013), they are regional in scale, typically  
477 conducted during a certain season, and are designed to sample a specific set of species or size-  
478 classes (e.g., large groundfishes). Different surveys also vary in timeframe, and availability of  
479 long time-series survey data is limited. On the other hand, survey data are available for a range  
480 of ecosystem types (from the tropics to high latitudes), thereby allowing the examination of  
481 model performance across ecological gradients.

#### 482 c. Species occurrence records

483 A major biological response to ocean changes is a shift in the distributions of marine species  
484 (Pinsky *et al.*, 2013; Poloczanska *et al.*, 2013), which can have further implications for marine  
485 ecosystems and LMR (Cheung *et al.*, 2010; Cheung *et al.*, 2013b). It is thus desirable for LMR  
486 models to realistically predict distributions for a wide range of species. A range of species  
487 distribution models have been applied to model LMRs under climate change (e.g., Jones and  
488 Cheung, 2015). The reliability of predicted species distributions are often examined using geo-  
489 referenced species occurrence records and test statistics, such as the Area Under Curve (AUC) of  
490 the Receiver Operating Characteristics (ROC). These records are collated from a range of  
491 sources including museum collections, scientific expeditions and surveys, and fisheries records.  
492 Many are now publicly accessible through databases, such as the Global Biodiversity  
493 Information Facility (GBIF) (Robertson *et al.*, 2014) and the Ocean Biodiversity Information  
494 System (OBIS) (Costello *et al.*, 2007), and have frequently been standardized for taxonomy and  
495 checked for quality. Species occurrence records have the advantage in having a much broader

496 spatial and taxonomic coverage than any single data source (e.g., from scientific survey only).  
497 However, problems with taxonomic misidentification, common names, synonyms, and errors in  
498 geo-referencing are still present. Confidence in species occurrence data may also be reduced due  
499 to sampling bias (Webb *et al.*, 2010). Specifically, information on locations where unsuccessful  
500 sampling has occurred is not always available, making it difficult to determine the areas where  
501 specific species are absent and therefore to interpret test statistics such as the AUC (Pearce and  
502 Boyce, 2006).

503  
504 To help inform the use of uncertain observational data in assessing model projections, a  
505 framework has been proposed to systematically assess the level of uncertainty associated with  
506 observational data particularly for climate change impact assessment (O'Connor *et al.*, 2015).  
507 This framework is based on evidence combined from theory, experiments and historical data  
508 with statistical analysis being undertaken to attribute any signals in observational data to climate  
509 change, thereby building confidence in the model. Such a framework will help identify cases  
510 where observational data are too uncertain to help assess model outputs e.g., with insufficient  
511 temporal and spatial coverage of observational data to reveal underlying trends and patterns.

512  
513 Post-processing of LMR model outputs is generally needed before they can be compared to  
514 empirical data, as there will inevitably be differences between LMR models due to variations in  
515 model structure and other factors. For example, output from species-based LMR models will be  
516 more directly comparable to empirical data. However, species-based LMR models may only  
517 include a subset of species or taxonomic groups that are included in the empirical data. In  
518 contrast, output from size-based models can easily be compared with aggregated LMR



519 production. However, the lack of explicit representation of taxonomic identity in size-based  
520 models makes their output difficult to compare to species- or population-specific data.  
521 Approximations can be made in some cases to convert information from size- or trophic- based  
522 models into taxonomic-based data. For example, the abundance and production of organisms at  
523 size > 1 m can be assumed to represent adult large pelagic fishes and can thus be compared to  
524 data from pelagic long-line catches. Functional group-based LMR models are intermediate  
525 between species-based and size-based models, and their outputs can be approximately converted  
526 to both taxonomic- or size-based aggregations. Thus, having identified the dominant taxonomic  
527 groups in a functional group, the dynamics of that functional group can be assumed to be  
528 representative of that taxonomic group. Functional groups that represent specific taxonomic  
529 groups of interest can also be included explicitly in the model (deYoung *et al.*, 2004; Griffith and  
530 Fulton, 2014).

531

### 532 **From quantifying uncertainty to assessing risk**

533 Given the large sources of uncertainty discussed in previous sections, a systematic exploration of  
534 potential future LMR states and the associated uncertainties is an important step towards a full  
535 risk assessment that would allow us to understand the potential impact of climate change on  
536 human societies through, for example, diminished food security, income or other ecosystem  
537 services. In general, risk consists of two components: (1) the magnitude of potential changes,  
538 and; (2) the probability of occurrence of such changes. Previous climatic risk assessments have  
539 involved both quantitative risk-based approaches and more qualitative, social vulnerability  
540 approaches (Dessai and Hulme, 2004), or a combination of both (Brown *et al.*, 2012).

541 Quantitative assessment generally involves identifying climate hazards and their probability of  
542 occurrence. For example, Li *et al.* (2009) assessed the drought risk for world crop production  
543 under climate change based on ensemble results from 20 GCM and six emission scenarios. The  
544 ensemble of projections was used to estimate probability density functions of drought disaster  
545 frequency. Their results show a consistent increase in drought risk in the middle and end of the  
546 21<sup>st</sup> century under climate change, leading to significant reductions in yield for major crops. In  
547 our case study of projecting changes in potential catches of sablefish in the Northeast Pacific  
548 (Figure 4), the probability of projecting a decrease in catch could be quantified by systematically  
549 exploring the envelope of uncertainty. Thorough estimates of risk can facilitate policy discussion  
550 for mitigation and/or adaptation in LMR management through the exploration of the potential for  
551 regrets/no-regrets policies and the associated costs and benefits (Polasky *et al.*, 2011). This  
552 approach to risk-based, ecosystem-based management has been developed for certain marine  
553 systems, for example in Australia (Hobday *et al.*, 2011). One area of risk assessment that  
554 remains particularly difficult to accurately quantify and yet important for guiding societal  
555 choices, is an understanding of “tail risk”, or risk from extreme and high-impact, but low-  
556 probability, events (Weitzman, 2011).

557

### 558 **Future direction of climate-LMR projections**

559 The many sources of uncertainty in climate-LMR projections and computational cost will always  
560 limit our ability to fully explore uncertainty in climate-LMR projections. However, the  
561 framework described here provides a basis for concerted effort to improve estimation of  
562 uncertainty ranges for climate-LMR projections and, eventually, reduce these ranges. As was the

563 case for physical climate projections, a climate-LMR ensemble offers a starting point. Systematic  
564 exploration of uncertainty space to identify prominent components for a given spatial scale, time  
565 horizon and variable of interest can guide research investment and accelerate progress toward  
566 more accurate estimates of uncertainty bounds. More rigorous and standardized comparison  
567 with observations (i.e., model metrics) must also play a central role in building confidence in  
568 projections. In combination, these steps should produce more robust risk estimates for policy  
569 formulation that will promote LMR sustainability in a changing climate.

570 While adoption of the framework described herein will improve climate-LMR projections,  
571 numerous challenges must still be overcome. Various unknowns pose a major challenge to  
572 exploring the real scope of uncertainties. Particularly, adaptive responses in nature to climate  
573 change, and by society to changes in LMRs, are difficult to predict and are poorly understood  
574 (Pinsky and Fogarty, 2012). There are also “unknown-unknowns”, such as ecological tipping  
575 points, which contribute to uncertainties and that cannot be assessed with our current knowledge.  
576 This problem could be partly addressed by developing scenarios that aim to explore the  
577 sensitivity of outputs to such uncertainties, such as a scenario incorporating high levels of  
578 biological and social adaptation. Additionally, when exploring structural uncertainty of the  
579 models, the sample of model structures is often assembled opportunistically based on existing  
580 models rather than strategically based on a systematic sampling of all plausible model structures.  
581 Furthermore, different climate-LMR models may not be entirely independent from one another  
582 as the models may be parameterized with similar datasets. This may result in biases in assessing  
583 the effects of model uncertainties on projections (Hawkins and Sutton, 2009). On the other hand,  
584 an ensemble of opportunities would be the most practical way to tackle the challenge of  
585 quantifying climate-LMR projection uncertainties and would help examine whether there is a

586 need for large-scale cooperative initiatives that provide substantial resources and facilities to  
587 address these challenges.

588 Observational data that are available for comparison with LMR models generally only cover a  
589 short period of time and a limited number of regions. This magnifies the issues regarding  
590 uncertainties associated with observation errors, making it more challenging to attribute the  
591 reasons for any discrepancies between observations and model predictions. Moreover, many  
592 LMR models use available observational data for parameterization, thus the scope of using  
593 additional data for model testing is limited. Careful selection of statistical and cross-validation  
594 techniques can help mitigate this problem (Arlot and Celisse, 2010). Further discussion and  
595 consensus amongst LMR modelers is needed to develop criteria to identify unrealistic models  
596 (i.e., what type and how many discrepancies are needed before a model is excluded from an  
597 ensemble). These challenges reiterate the need to improve the sharing of observational data  
598 between scientists, institutes and countries and develop data facilities to support their use in  
599 testing climate-LMR projections (Hollowed *et al.*, 2013).

600 Scenario development has not matured for LMR assessment. Scenarios specifically tailored for  
601 marine-related sectors are very limited, while existing assessments adopt scenarios that are used  
602 for more general purposes (Millennium Ecosystem Assessment, 2005). These scenarios may not  
603 account for key uncertainties in the projected pathways of LMRs. In relation to this, fisheries  
604 models linking fishing to changes in LMRs, and the socio-economic conditions that are used to  
605 generate LMR scenarios are only starting to be developed for global- and basin- scale LMRs,  
606 although much effort has focused on regional- and local- scale fishing fleet dynamic models (van  
607 Putten *et al.*, 2012) and management strategy evaluation (MSE) models. All existing global- or  
608 basin-scale LMR models either do not have explicit fisheries components or have simple

609 assumptions of stock- or region- specific fishing mortality rates. Only recently has a global scale  
610 LMR model study included a spatially-explicit fishing dynamics model to simulate changes in  
611 fishing effort (Christensen *et al.*, 2015). However, there is a need to improve efforts such as this  
612 to develop additional LMR-specific scenarios representing human activities before meaningful  
613 comparison of scenario uncertainties can be undertaken.

614 Understanding where uncertainty comes from and how it interacts with model components is  
615 necessary to improve the interpretation of model projections and to inform policy. Improving  
616 the quantification of uncertainties will therefore be a major area of development in climate-LMR  
617 projections to inform global and regional assessments of climate change impacts, vulnerability  
618 and adaptation on marine ecosystems and related sectors.

619

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627

628

629 Table 1. Summary of different types of uncertainties in LMR models.

Types of uncertainties	Description	Examples	
		<b>Coupled</b>	<b>atmospheric, Fish and fisheries models ocean and biogeochemical models</b>
<b>Internal variability</b>	Natural variations of physical, biogeochemical and ecological processes that contribute random variability to projections of LMRs	El Niño-Southern Oscillation (ENSO) and Pacific Decadal Oscillation (PDO), North Atlantic Oscillation (NAO), Atlantic Multidecadal Oscillation (AMO), locations of gyre boundaries	Predator-prey dynamics, spatial and temporal variations in fish populations not arising from deterministically modeled climate change signal.
<b>Model uncertainty</b>	a. Parameter Specific parameter values used in the formulae determining the behavior of the models	Parameters controlling sub-grid scale oceanographic processes, phytoplankton growth, zooplankton grazing, biogeochemical transformations, and detritus remineralization.	Values of the parameter describing diet composition, dispersal rate, production and consumption rates, trophic interactions and other ecological/anthropogenic processes represented in the models. If variations in parameter values reflect an alteration of model architecture or design, it

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should belong to the structural uncertainty category.

b. Structure      Differences in abstraction, understanding and representation of the system through different model architecture, design and assumptions, the method of representing space/time, and the kinds of ecological processes, human and natural drivers included.      Size-based vs functional group based approaches, grid resolution; the number of nutrients included in biogeochemical models; the number of functional groups included in a lower trophic level or fish model; differences in representing food web structure, fish movement, and different life history stages in fish models.

**Scenario uncertainty**      Differences in the natural and/or anthropogenic forcing that drive the model simulation      Representative Concentration Pathways (RCPs)      Shared Socio-economic Pathways (SSPs), spatial and temporal changes in fisheries, and how all of these influence model components directly or indirectly

631 Table 2. Selected case studies that explored different aspects of uncertainties in projections of  
 632 aquatic (marine and freshwater) biological resources under climate change.

<b>Spatial scale</b>	<b>Selected case studies</b>	<b>Explored uncertainties and conclusions</b>
Global	Variability of projections of distribution and patterns of species turnover across three different species distribution models for over 800 commercially exploited fishes and invertebrates in the world under two greenhouse gas emission scenarios (Jones and Cheung, 2015).	<ul style="list-style-type: none"> <li>- Structural uncertainties of species distribution models;</li> <li>- Scenario uncertainties of greenhouse gas emission pathways.</li> <li>- Larger variability in projections exists between greenhouse gas emission scenarios (RCP2.6 and RCP8.5) than between three different species distribution models.</li> </ul>
Regional (UK waters)	Projecting changes in maximum catch potential and profitability from fishing 31 key commercially targeted fish species primarily inhabiting UK waters using different climate models, species distribution modelling approaches and socio-economic scenarios (Jones <i>et al.</i> , 2014). Three fisheries and socio-economic scenarios were designed based on key	<ul style="list-style-type: none"> <li>- Structural uncertainties of species distribution models and climate models;</li> <li>- Scenario uncertainties of greenhouse gas emission, fisheries and socio-economics pathways;</li> <li>- Scenario (climate, fisheries and socio-economic) uncertainty</li> </ul>



variables identified in the Alternative Future Scenario for Marine Ecosystems (AFMEC) scenarios.

dominates over structural uncertainty of climate and biological models.

Regional (Central North Pacific Ocean) Uncertainty of a trophodynamic model (Ecopath with Ecosim) was explored using Monte Carlo simulation. Confidence limits of key input parameters were set based on the reliability of the data, as indicated by the data type. Results from 500 dynamic simulations (each involving up to several thousand iterations to find a balanced model) were used to construct 95 % confidence intervals for the derived biomass time series (Kearney *et al.*, 2012).

- Parameter uncertainty of the ecological models.

Regional (Eastern U.S. coast) Using experimentally-derived thermal tolerance limits to project range shift of gray snapper (*Lutjanus griseus*) in estuaries along eastern US coast. Projections were driven by temperature simulated from 23 different climate models, two thermal tolerance metrics under three different emission scenarios (Hare *et al.*, 2012).

- Parameter uncertainty of range shift model;  
- Structural uncertainties of climate models;  
- Scenario uncertainties of greenhouse gas emission pathways.  
- Different species distribution models contributed the largest

variation in projections, followed by different General Circulation Models (GCMs). The contribution of variability from different GCMs increased over time and to a level that is comparable to variability from different species distribution models for end of 21<sup>st</sup> century projections. Different observation datasets had a small influence on the overall variability of the projections.

<p>Regional (freshwater ecosystems in France)</p>	<p>Projection of distribution shifts of 35 species of freshwater fishes in France across 100 random subsets of observation data, seven species distribution models and climate projections from 12 climate models, resulting in 8400 different potential futures projections (Buisson <i>et al.</i>, 2010).</p>	<ul style="list-style-type: none"> <li>- Parameter uncertainty of species distribution models;</li> <li>- Structural uncertainties of species distribution models and climate models;</li> <li>- Scenario uncertainty of greenhouse gas emission pathways.</li> <li>- Uncertainty about thermal limits of the species dominates over model or scenario uncertainties</li> </ul>
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634 Table 3. Examples of observation datasets and the model metrics for comparison with LMR  
 635 model outputs.

<b>Data type/ Model metric</b>	<b>Timeframe</b>	<b>Spatial aggregation</b>	<b>Taxonomic resolution</b>	<b>Examples of data sources</b>
<b>Fisheries data</b>				
Total fisheries catch potential	Average from 1950 - 2010	Global and large marine ecosystems	Aggregated	Sea Around Us project
Species composition of catch (or the mean temperature of catch)	Annual from 1950 - 2010	Global and large marine ecosystems	Exploited taxa	Sea Around Us project
CPUE of large pelagic fishes	1970s to 2000s	Global	Large pelagic fishes (tunas)	Regional Fisheries Management Organizations for tunas and billfishes e.g., Myers and Worm (2003)
<b>Survey data</b>				
Rate of range shift of marine species	Average from 1970s to 2000s	Regional (North America continental shelf, North Sea)	By species of fishes and invertebrates	Pinsky <i>et al.</i> (2013), ICES's International Bottom Trawl Survey
Community composition	1960s to 2010s	Regional (continental)	By species of	Worm <i>et al.</i> (2009)

		shelves around the world)	fishes and invertebrates vulnerable to bottom trawls	
Variability in abundance driven by large-scale oceanographic changes	1951 – Present	Regional: California Current	Larvae and eggs of exploited and unexploited fishes	CalCOFI (Moser <i>et al.</i> , 2001)
	1931 - Present	North Sea	Exploited and unexploited fishes	Continuous Plankton Recorder (CPR) Survey
	1970s – 2000s	North America continental shelf	By species of fishes and invertebrates	Pinsky <i>et al.</i> (2013)
<b>Occurrence record</b>				
Occurrence of marine species	Mostly since the 20 <sup>th</sup> century	Global	All marine taxa	OBIS (Costello <i>et al.</i> , 2007)

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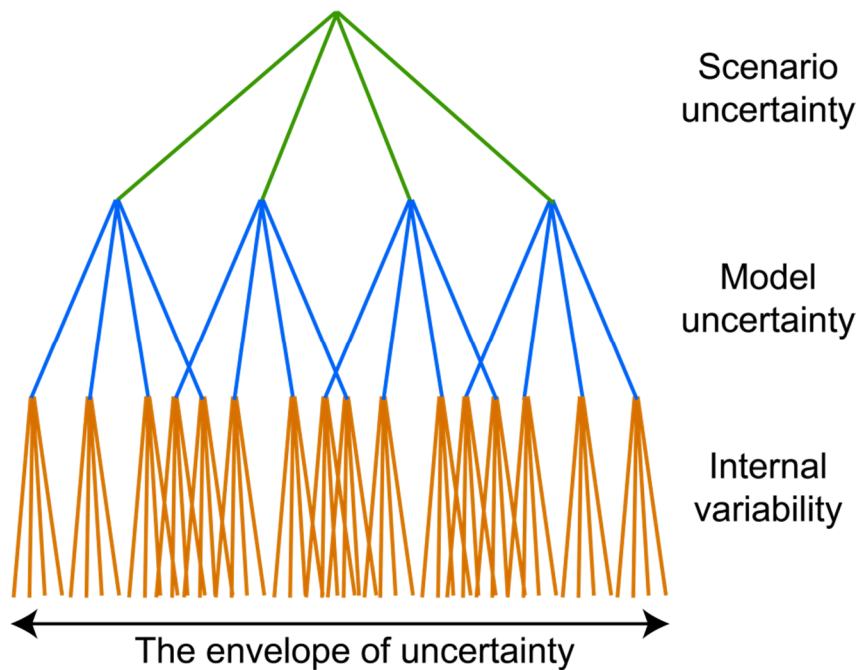


Figure 1. Schematic example illustrating cascades of uncertainties in LMR projection (modified from Wilby and Dessai (2010)). For a particular time horizon and spatial scale, the range, or envelope, of possible outcomes includes contributions from scenario uncertainty (green), model uncertainty (blue) and internal variability (orange). The cascades of uncertainties apply to each of the sub-components of climate-LMR models. Uncertainties from each model sub-component may be additive or multiplicative. In this schematic diagram, the width of each uncertainty level does not imply the magnitude of the uncertainty. E.g., internal variability may be larger than scenario uncertainty and *vice versa*.

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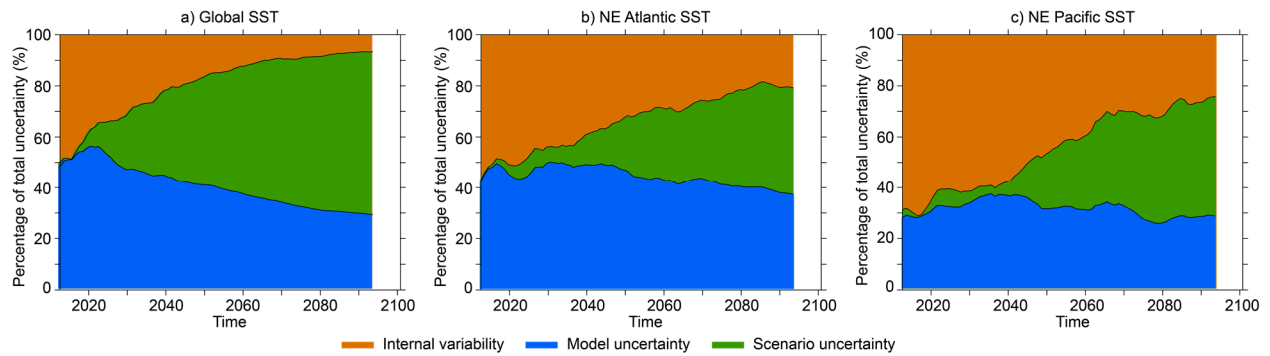


Figure 2. The relative importance of each source of uncertainty in annual mean sea surface temperature projection is shown by the fractional uncertainty for (a) global mean, (b) northeast Atlantic, and (c) northeast Pacific in the 21<sup>st</sup> century. Uncertainties are separated into three components: internal variability (orange), model uncertainty (blue), and scenario uncertainty (green). The percentage of total uncertainty is calculated from dividing the level of uncertainty from the specific component by the sum of the three types of uncertainties. For internal variability, the standard deviation of annual mean SST from the GFDL ESM2M ensemble is calculated year-by-year. The same procedure has been applied for model uncertainty, but a 10-yr running mean (longer than the typical ENSO period) is first applied to the individual CMIP5 model projections.

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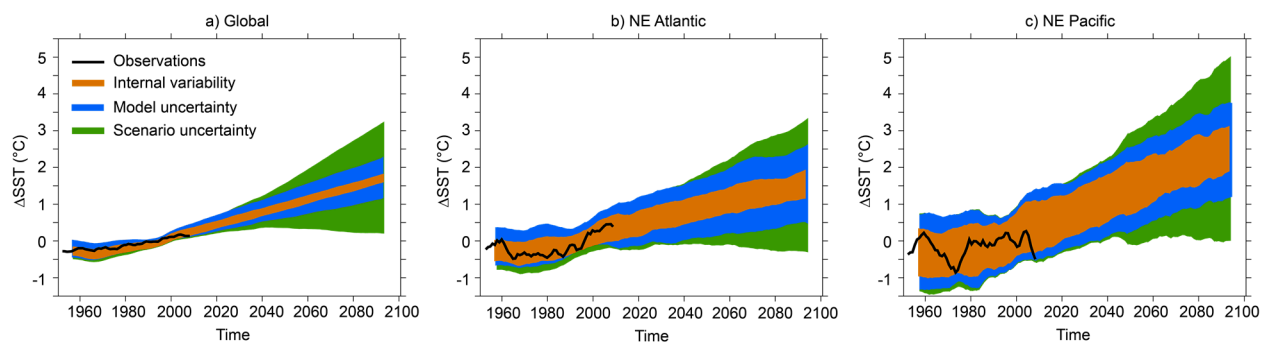
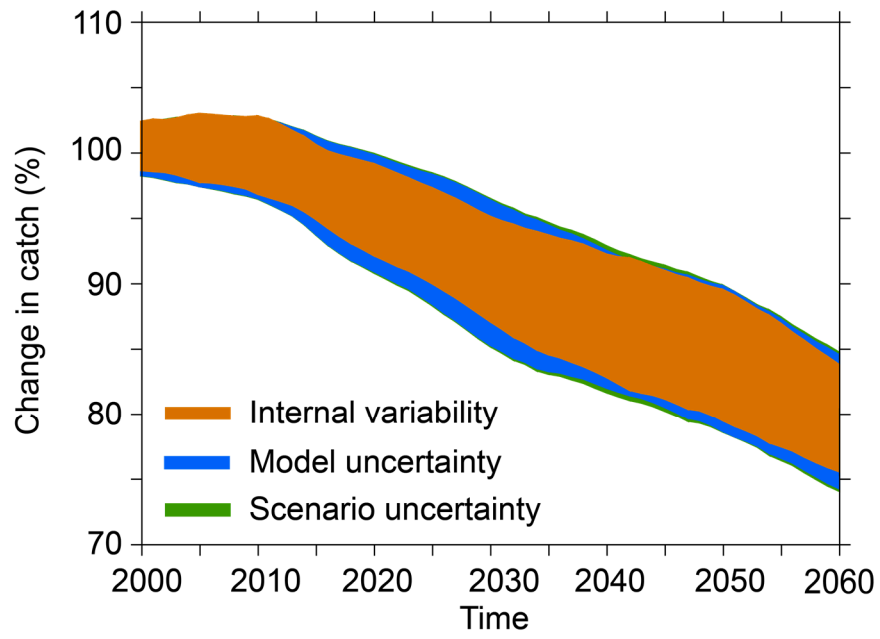


Figure 3. Changes in annual average sea surface temperature (10-year running mean) for (a) global mean, (b) northeast Atlantic, and (c) northeast Pacific relative to the 1986-2005 mean. SST observations (black line) are based on Smith *et al.* (2008). The uncertainty area was calculated by adding and subtracting the errors from each uncertainty source (internal variability: orange, model uncertainty: blue, scenario uncertainty: green) to and from the ensemble-mean projection of 15 CMIP5 models. Errors from different uncertainty sources are assumed to be additive.

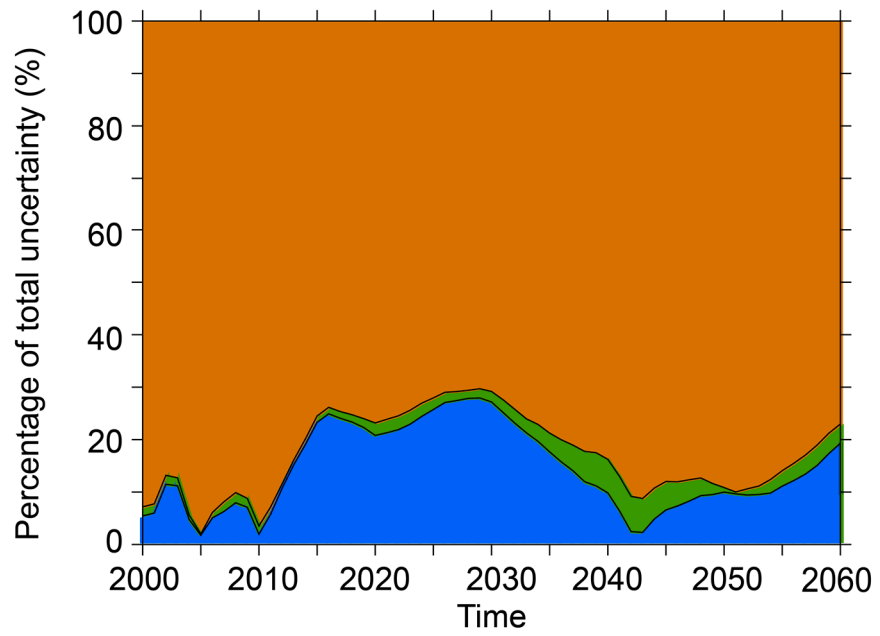
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A



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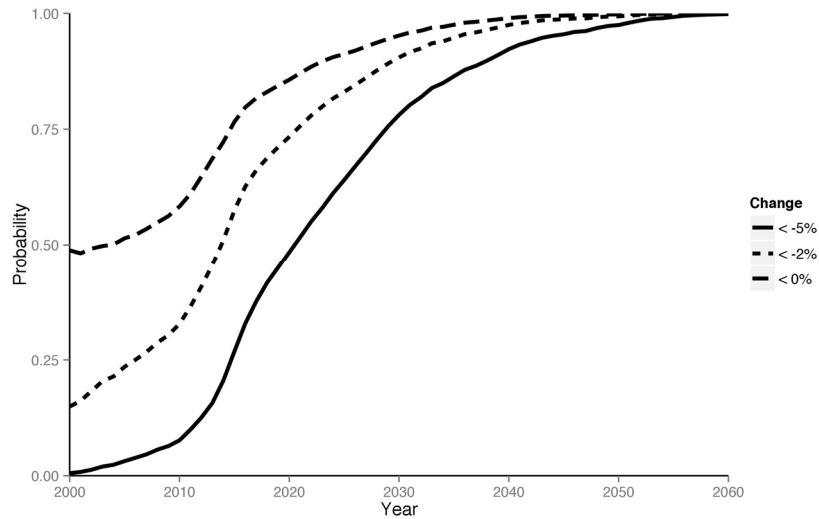


Figure 4. Projected changes in maximum potential catches of *Anoplopoma fimbria* from 2000 to 2060 under climate change. The projections were generated from using three versions of Dynamic Bioclimate Envelope Models (DBEM) (Cheung et al. under review), driven by outputs from GFDLESM2M. Internal variability was estimated from projected changes in catch potential driven by outputs from 20 ensemble members of GFDL ESM2M (Rodgers *et al.* 2015). (A) Projected changes in maximum potential catch and their standard deviation resulting from the three different types of uncertainties. (B) the relative contribution of each type of uncertainty, expressed as the proportion of total uncertainty, and (C) the probability of projecting a decrease in catch potential of more than 0% (dashed line), 2% (dotted line) and 5% (solid line). Model uncertainty represents variation of projections from the three versions of DBEM. Scenario uncertainty represents variations in projections between RCP2.6 and RCP8.5.

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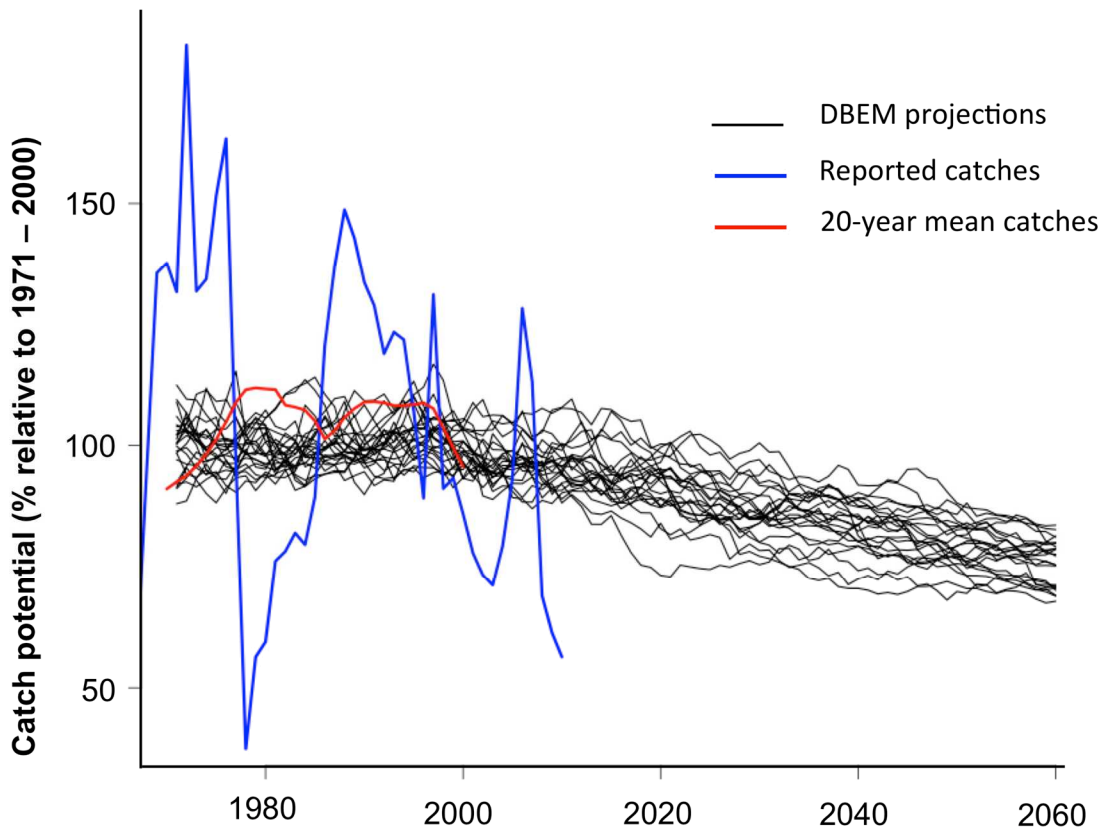
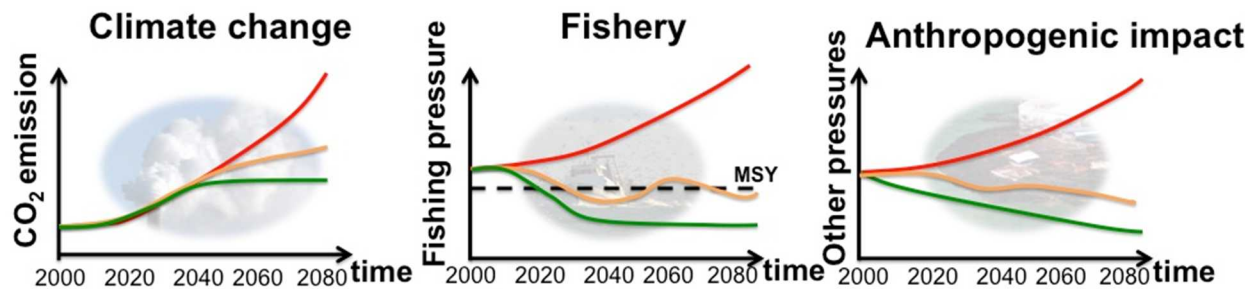


Figure 5. Comparison between projected changes in annual mean catch potential (relative to 1971 – 2000) using the three versions of DBEM and 20 GFDL ESM2M ensemble members under the RCP8.5 scenario (grey lines) with the reported catches (from Sea Around Us: [www.seaaroundus.org](http://www.seaaroundus.org)) of sablefish in the Northeast Pacific (Gulf of Alaska Large Marine Ecosystem) (blue line). Reported catches are also smoothed by a 20-year running mean (red line).

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651

652 Figure 6. Schematic diagram showing an example of potential standardized sets of scenarios to  
653 be developed to explore scenario uncertainties. The red, yellow and green lines represent  
654 different scenario pathways to be explored by climate-LMR models. Anthropogenic impacts may  
655 include contaminant level, invasive species and habitat change.

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